

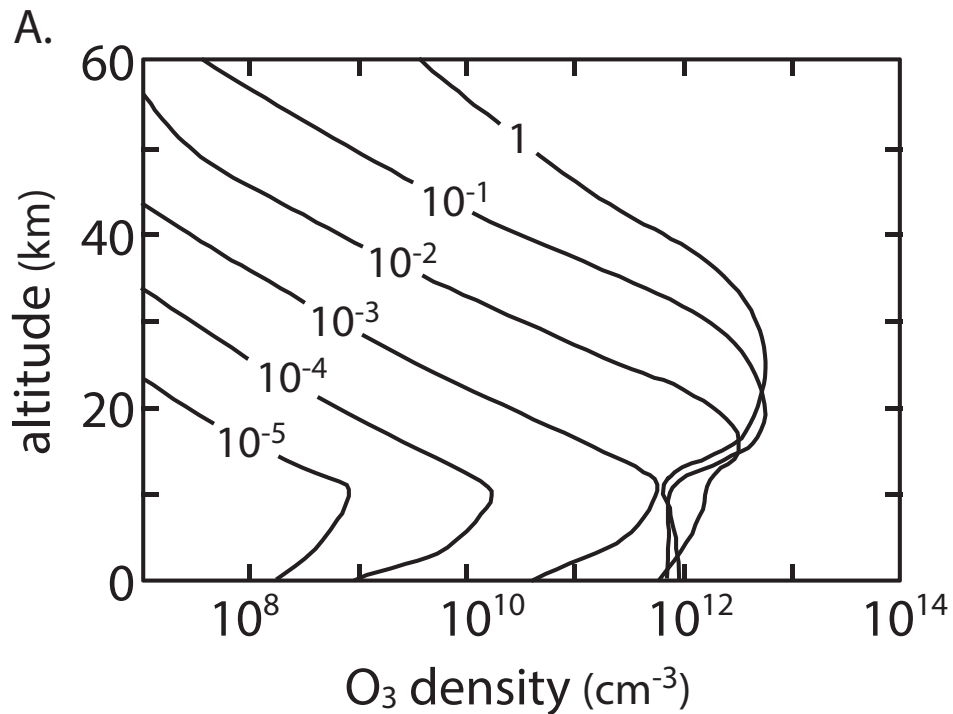
methane greenhouse?

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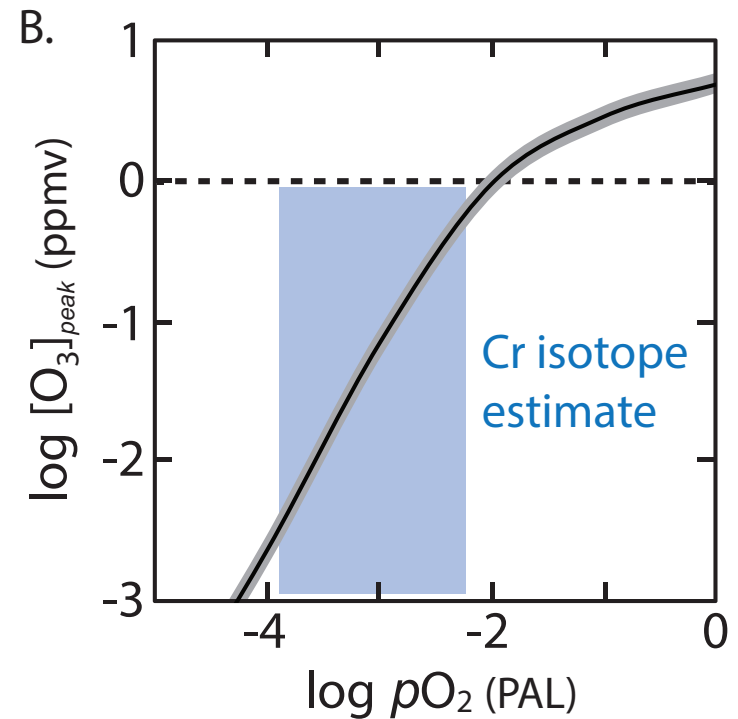
oxygen (O_2) & ozone (O_3)

less O_2 = less O_3

O_3 is abiotically produced from O_2 in the atmosphere



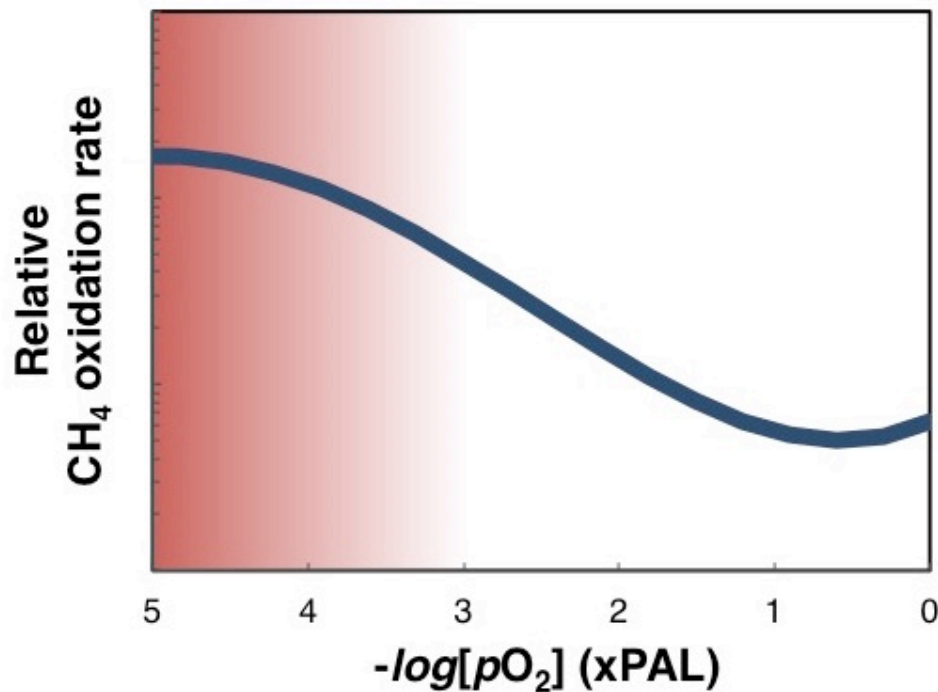
*after Kasting & Donahue [1980]



Reinhard, Olson, Lyons, Kasting [*In prep*]

oxygen (O₂) & methane (CH₄)

Ozone protects CH₄ (and other biosignature gases)

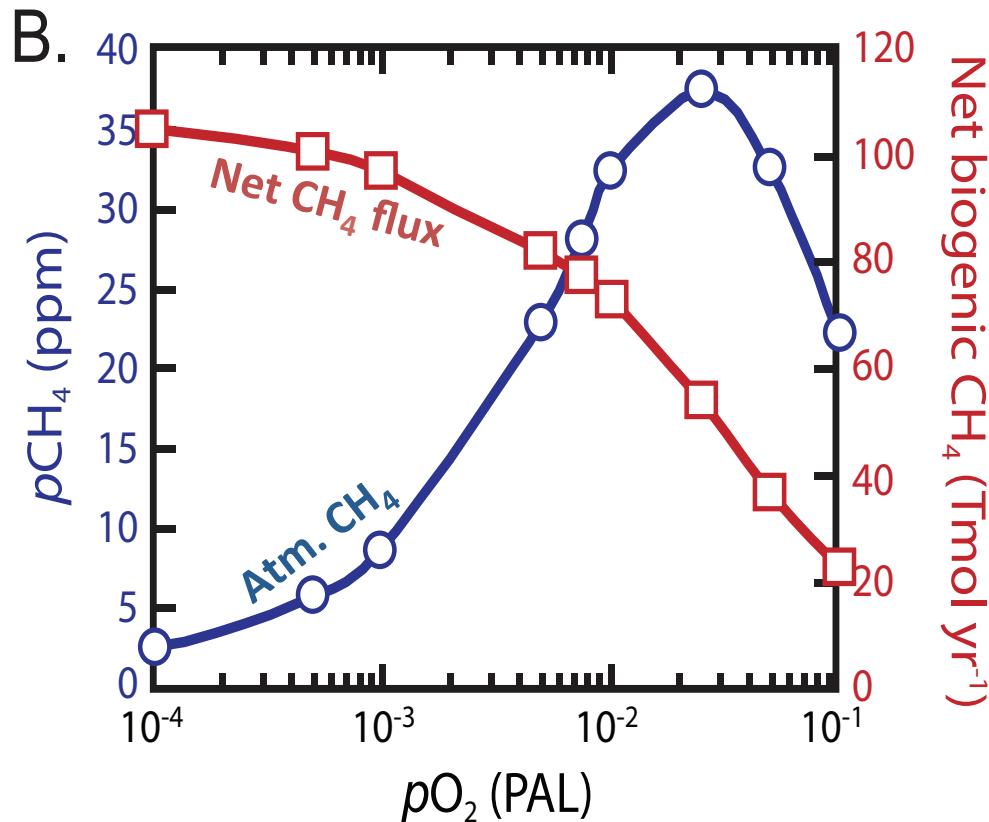


*after Goldblatt et al. [2006]

- CH₄ destruction is muted when O₃ layer is well developed.
- Proterozoic atmosphere was **more oxidizing** toward CH₄ than the present O₂ rich atmosphere

oxygen (O₂) & methane (CH₄)

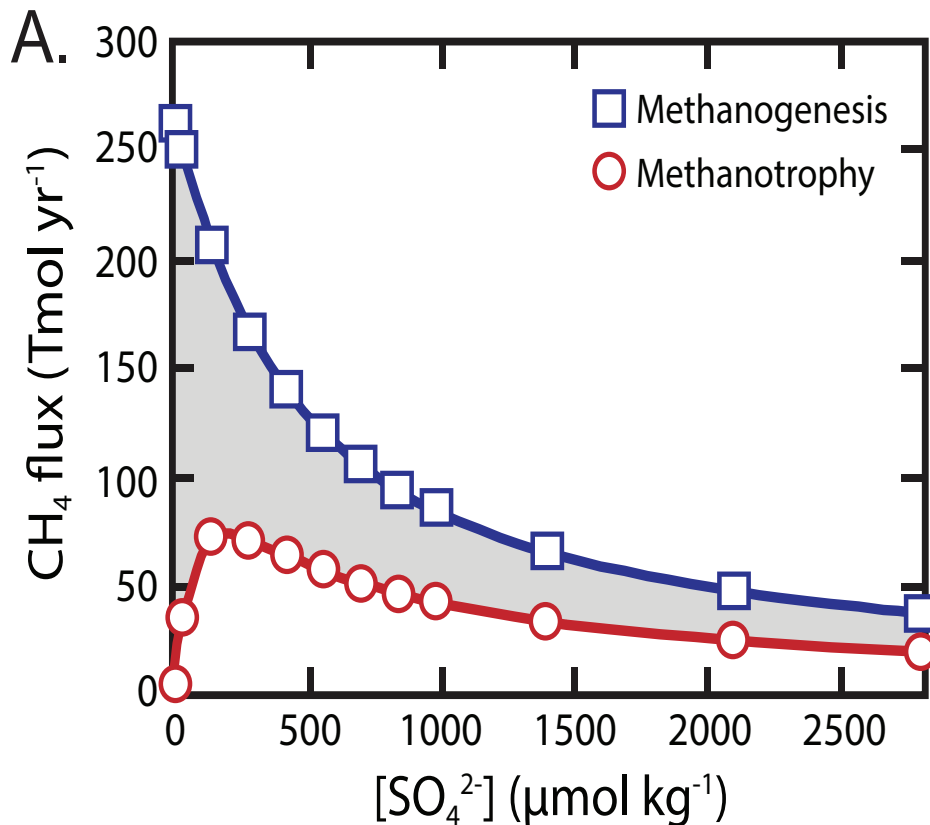
CH₄ production declines with O₂—but ozone favors CH₄ accumulation



- Maximum pCH_4 occurs for peak preservation—not peak production.
- **Elevated pCH_4 requires relatively high pO_2 (O₃).**

sulfate (SO_4) & methane (CH_4)

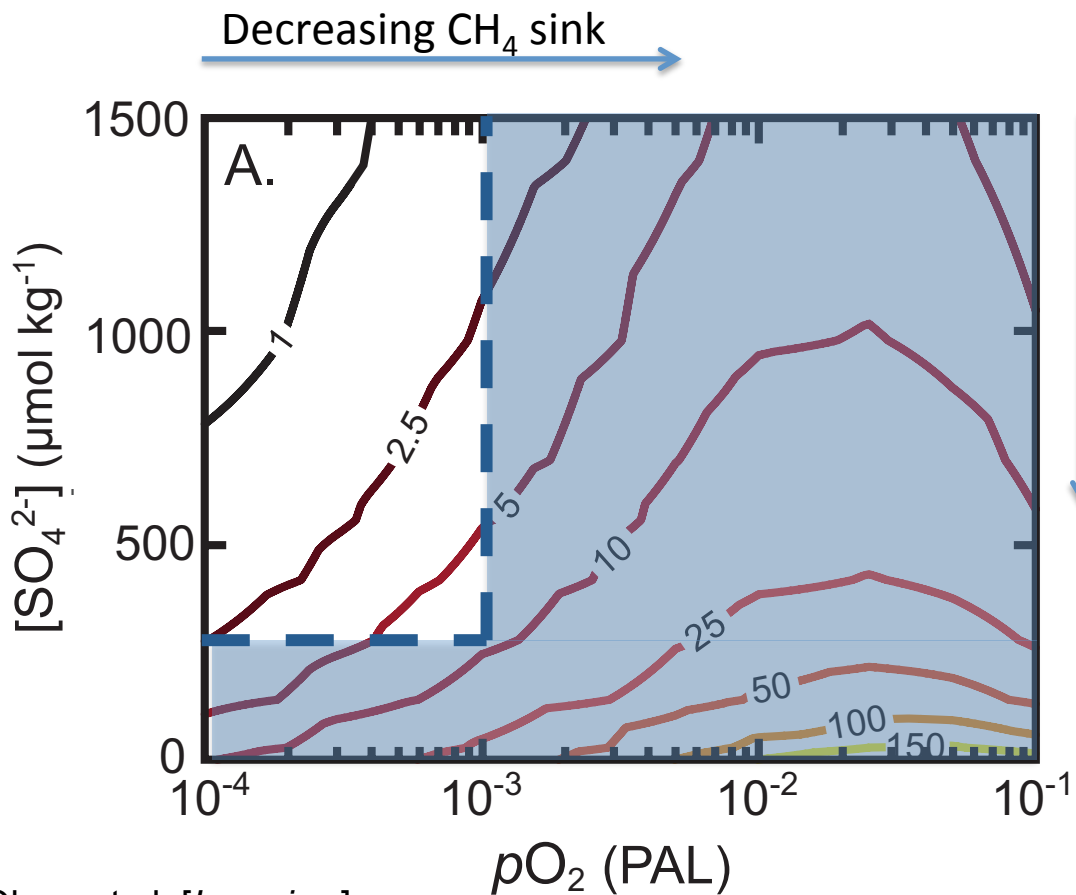
Oceanic SO_4 is the primary control on biogenic CH_4 fluxes



- Both CH_4 production and CH_4 preservation are disfavored by SO_4
- **Substantial CH_4 supply to the atmosphere requires exceptionally low SO_4 concentrations**

methane (CH₄)

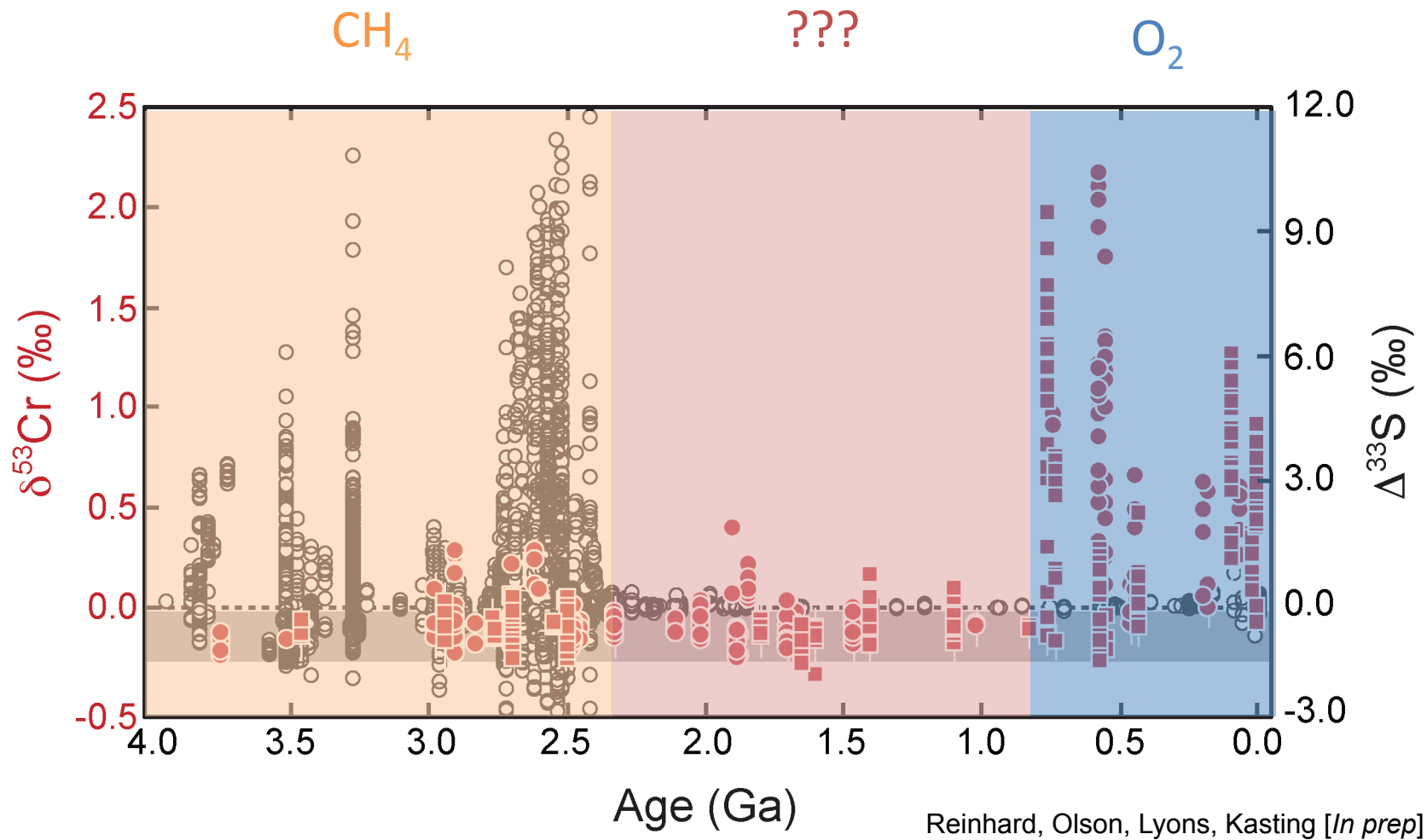
mid-Proterozoic pCH₄ was probably similar to today



Olson et al. [In review]

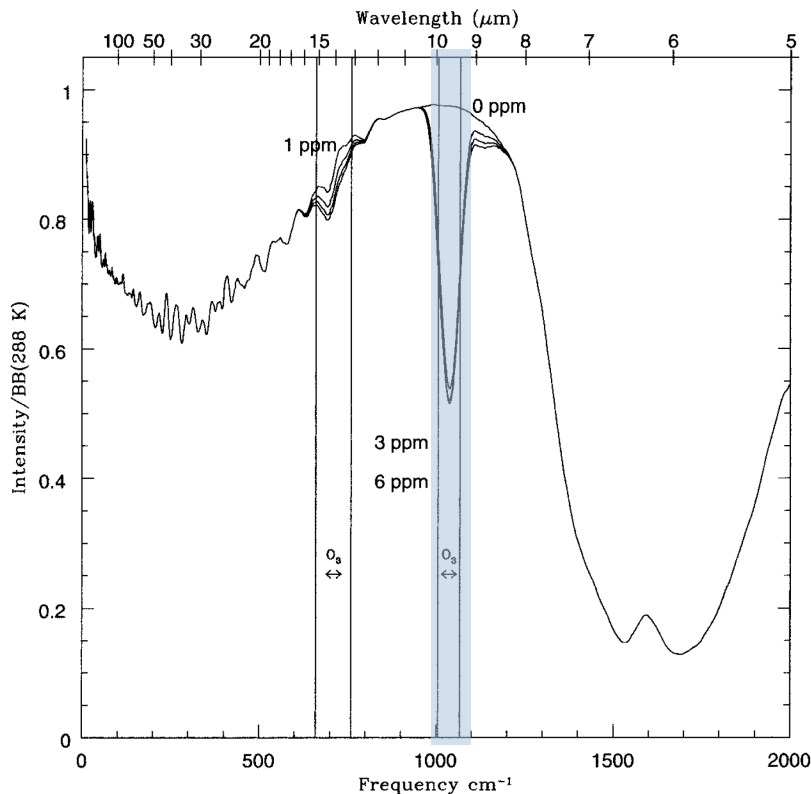
- The composition of the Proterozoic greenhouse remains unknown.
- **CH₄ is not a viable biosignature for the second half of Earth history**

biosignature gases through time



ozone (O₃)

O₃ (abiotic) can be a proxy for trace biogenic O₂



Des Marais et al. [2002]

- O₃ has strong absorption features
- At the upper range of reconstructed mid-Proterozoic pO_2 values detection of O₃ may be possible
- **O₃ (abiotic) vulnerable to false positives**